

Flow and Reactive transport in shallow aquifer

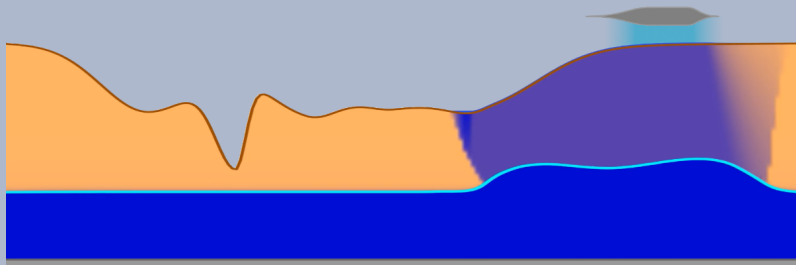
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24/06/2024

Journées transport réactif, LMPA Calais

Objective and outline

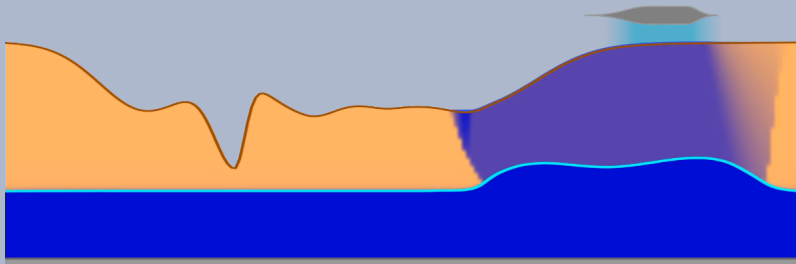


Context

Describe **efficiently** pollutant transport in **shallow aquifers**.

- ↔ **Typically** nitrates pollution of the watertable due to agriculture
- ↔ **Modelling**: Flow + Transport + Chemical reaction

Objective and outline



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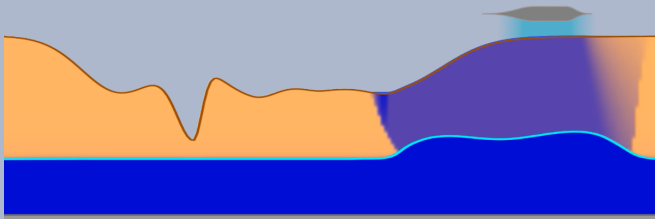
Dificulties in the flow description

- Flow in porous media: **3d-Richards model**
 - ↔ **non-linear + 3d = Difficult** (mathematically and numerically)
- **Lagre geometry**
 - ↔ **long time scale seems important**
- Chemical reactions may occur **above the water table**
 - ↔ **short time scale seems important**

Objective: Propose simplified models to describe the **flow** and the **reactive transport**

Strategy: use the shallow geometry to simplify the flow modelling

Outline

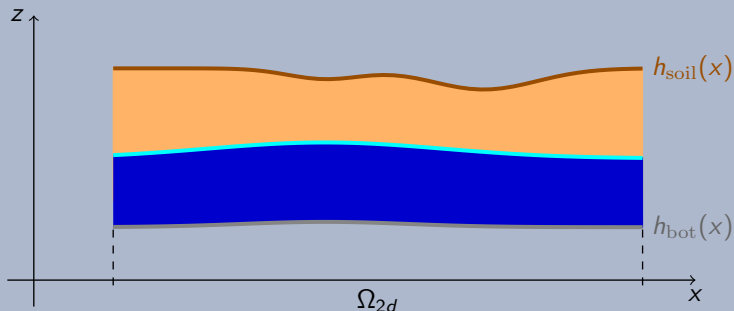


- 1 Geometry and 3d-Richards equations
- 2 Dominant components of the flow: effective problems
- 3 A coupled model coupling dominant components
- 4 Numerical scheme
- 5 Considering reactive transport
- 6 Reactive transport: Dominant component and coupled model
- 7 Numerical scheme

Plan

- 1 **Geometry and 3d-Richards equations**
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Notations and Geometry



- Ω_{2d} : horizontal domain
- x : horizontal variable
- z : vertical variable
- e_3 : unitary vertical vector
- $h_{\text{bot}}(x)$: level of the bottom
- $h_{\text{soil}}(x)$: level of the soil

$$\Omega_{3d} := \left\{ (x, z) \mid x \in \Omega_{2d}, z \in]h_{\text{bot}}(x), h_{\text{soil}}(x)[\right\}$$

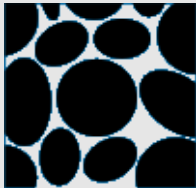
Physical quantities and Richards equation

- P : Pressure head
- H : Hydraulic head
- \mathbf{v}^f : Fluid velocity

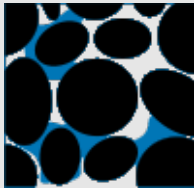
- θ : Water saturation
- k : Hydraulic conductivity

Richards Equation in $]0, T[\times \Omega_{3d}$

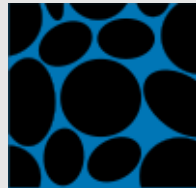
$$\begin{cases} \frac{\partial \theta}{\partial t} + \operatorname{div} \mathbf{v}^f = 0 \\ \mathbf{v}^f = -k \nabla H \\ H = P + z \end{cases}$$



$$\theta \simeq 0$$



$$0 < \theta < 1$$



$$\theta = 1$$

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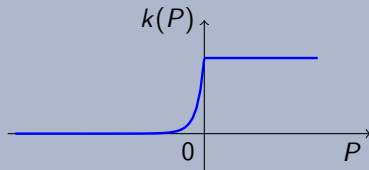
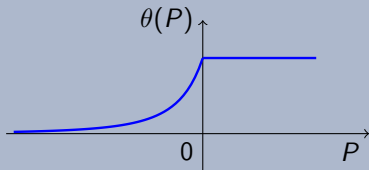
- $\theta(P)$: Water saturation
- $k(P)$: Hydraulic conductivity

Richards Equation in $]0, T[\times \Omega_{3d}$

$$\begin{cases} \frac{\partial \theta(P)}{\partial t} + \operatorname{div} \mathbf{v}^f = 0 \\ \mathbf{v}^f = -k(P) \nabla H \\ H = P + z \end{cases}$$

Hypothesis:

- $\theta = \theta(P)$, $k = k(P)$
- $\theta(P) = 1 \iff P \geq 0$

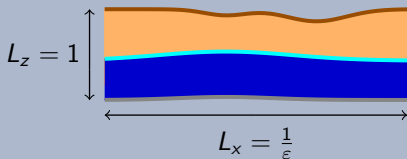


Plan

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Dominant components of the flow

- **Objective:**
exhibits **dominant component** of the flow when the aquifer is very large
- **Strategy:**
 - Aquifer large and shallow: $L_x = \frac{1}{\varepsilon}$, $L_z = 1$ then make $\varepsilon \rightarrow 0$



Physical variables (x, z)

Dominant components of the flow

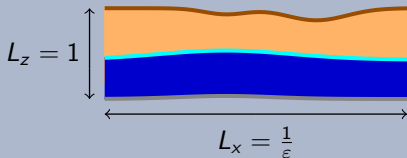
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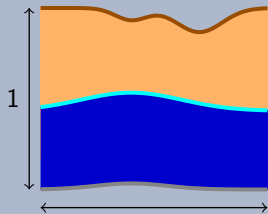
- **Strategy:**

- Aquifer large and shallow: $L_x = \frac{1}{\varepsilon}$, $L_z = 1$ then make $\varepsilon \rightarrow 0$
- Change of variable

$$\bar{x} = \frac{x}{L_x} = \varepsilon x, \quad \bar{z} = \frac{z}{L_z} = z$$



Physical variables (x, z)



Dimensionless variables (\bar{x}, \bar{z})

Rescaled Richards equation

Richards Equation

$$\frac{\partial \theta(P)}{\partial t} - \operatorname{div}_x (k(P) \nabla_x H) - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) = 0$$

Rescaled Richards equation

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$$\frac{\partial \theta(P)}{\partial t} - \operatorname{div}_x (k(P) \nabla_x H) - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) = 0$$

- $L_x = \frac{1}{\varepsilon}$ characteristic size of Ω_{2d} $\Rightarrow \bar{x} = \frac{x}{L_x}$
- $L_z = 1$ characteristic vertical size of Ω_{3d} $\Rightarrow \bar{z} = \frac{z}{L_z}$
- T characteristic time $\Rightarrow \bar{t} = \frac{t}{T}$

Rescaled Richards equation

Richards Equation

$$\frac{\partial \theta(P)}{\partial t} - \operatorname{div}_x (k(P) \nabla_x H) - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) = 0$$

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Dimensionless Richards Equation

$$\frac{1}{T} \frac{\partial \theta(\bar{P})}{\partial \bar{t}} - \varepsilon^2 \operatorname{div}_{\bar{x}} (k(\bar{P}) \nabla_{\bar{x}} \bar{H}) - \frac{\partial}{\partial \bar{z}} \left(k(\bar{P}) \frac{\partial \bar{H}}{\partial \bar{z}} \right) = 0$$

What for the time ?

$$\frac{1}{T} \frac{\partial \theta(\bar{P})}{\partial \bar{t}} - \varepsilon^2 \operatorname{div}_{\bar{x}} \left(k(\bar{P}) \nabla_{\bar{x}} \bar{H} \right) - \frac{\partial}{\partial \bar{z}} \left(k(\bar{P}) \frac{\partial \bar{H}}{\partial \bar{z}} \right) = 0$$

Time consideration

- Short time scale: $T = 1$

$$\frac{\partial \theta(\bar{P})}{\partial \bar{t}} - \varepsilon^2 \operatorname{div}_{\bar{x}} \left(k(\bar{P}) \nabla_{\bar{x}} \bar{H} \right) - \frac{\partial}{\partial \bar{z}} \left(k(\bar{P}) \frac{\partial \bar{H}}{\partial \bar{z}} \right) = 0$$

- Long time scale: $T = \frac{1}{\varepsilon^2}$

$$\varepsilon^2 \frac{\partial \theta(\bar{P})}{\partial \bar{t}} - \varepsilon^2 \operatorname{div}_{\bar{x}} \left(k(\bar{P}) \nabla_{\bar{x}} \bar{H} \right) - \frac{\partial}{\partial \bar{z}} \left(k(\bar{P}) \frac{\partial \bar{H}}{\partial \bar{z}} \right) = 0$$

Effective problems: short-time scale

Asymptotic expansion

Formally:

$$\bar{P}_\varepsilon = \bar{P}_0 + \varepsilon \bar{P}_1 + \varepsilon^2 \bar{P}_2 + \dots \quad \bar{H}_\varepsilon = \bar{H}_0 + \varepsilon \bar{H}_1 + \varepsilon^2 \bar{H}_2 + \dots$$

$$\frac{\partial \theta(\bar{P}_\varepsilon)}{\partial t} - \varepsilon^2 \operatorname{div}_{\bar{x}} \left(\bar{k}(\bar{P}_\varepsilon) \nabla_{\bar{x}} \bar{H}_\varepsilon \right) - \frac{\partial}{\partial \bar{z}} \left(k(P_\varepsilon) \frac{\partial \bar{H}_\varepsilon}{\partial \bar{z}} \right) = 0$$

1d-Richards problem

$$\phi \frac{\partial \theta(\bar{P}_0)}{\partial \bar{t}} - \frac{\partial}{\partial \bar{z}} \left(\bar{k}(\bar{P}_0) \frac{\partial \bar{H}_0}{\partial \bar{z}} \right) = 0 \quad \text{in } \Omega_{3d}$$

Effective problems: short-time scale

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No time for the water to have a significant horizontal displacement.

Effective problems: long-time scale

$$\varepsilon^2 \frac{\partial \theta(\bar{P}_\varepsilon)}{\partial t} - \varepsilon^2 \operatorname{div}_{\bar{x}} \left(\bar{k}(\bar{P}_\varepsilon) \nabla_{\bar{x}} \bar{H}_\varepsilon \right) - \frac{\partial}{\partial \bar{z}} \left(\bar{k}(\bar{P}_\varepsilon) \frac{\partial \bar{H}_\varepsilon}{\partial \bar{z}} \right) = 0$$

Dupuit-like problem

$$\left\{ \begin{array}{l} \bar{H}_0(t, x, z) = \bar{H}_0(t, x) \\ \end{array} \right.$$

Effective problems: long-time scale

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Dupuit-like problem

$$\begin{cases} \bar{H}_0(t, x, z) = \bar{H}_0(t, x) \\ \frac{\partial}{\partial \bar{t}} \int_{\bar{h}_{\text{bot}}}^{\bar{h}_{\text{soil}}} \theta(\bar{P}_0) d\bar{z} - \operatorname{div}_{\bar{x}} \left(\tilde{k}(\bar{H}_0) \nabla \bar{H}_0 \right) = 0 \\ \tilde{k}(\bar{H}_0) = \int_{\bar{h}_{\text{bot}}}^{\bar{h}_{\text{soil}}} \bar{k}(\bar{P}_0) d\bar{z} \end{cases}$$

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Vertical flow seems instantaneous: stationary state is reached.

Conclusion of asymptotic analysis

Short-time scale: 1d vertical Richards problem

$$\frac{\partial \theta(\bar{P}_0)}{\partial \bar{t}} - \frac{\partial}{\partial \bar{z}} \left(\bar{k}(\bar{P}_0) \frac{\partial \bar{H}_0}{\partial \bar{z}} \right) = 0$$

Long-time scale: 2d horizontal “Dupuit’s problem”

$$\begin{cases} \bar{H}_0(\bar{t}, \bar{x}, \bar{z}) = \bar{H}_0(\bar{t}, \bar{x}) \\ \frac{\partial}{\partial \bar{t}} \int_{\bar{h}_{\text{bot}}}^{\bar{h}_{\text{soil}}} \theta(\bar{P}_0) d\bar{z} - \text{div}_{\bar{x}} \left(\tilde{k}(\bar{H}_0) \nabla_{\bar{x}} \bar{H}_0 \right) = 0 \end{cases}$$

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Conclusion: 3d-Richards problem \rightarrow effective equations

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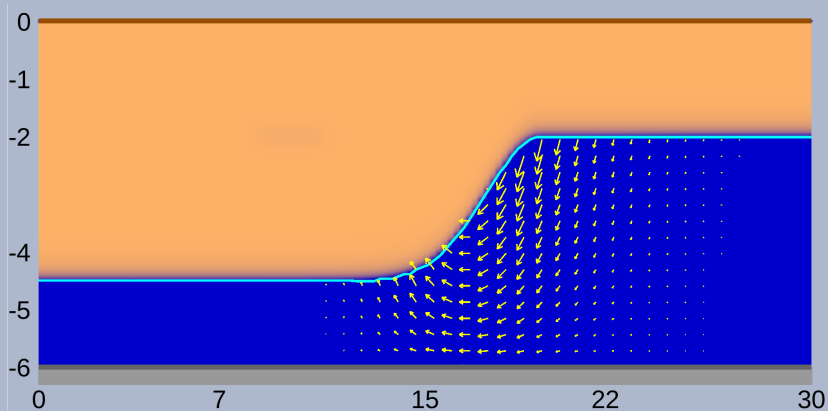
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Conclusion: 3d-Richards problem \rightarrow effective equations

Strategy: Simpler coupled problem \rightarrow effective equations

Numerical experiment

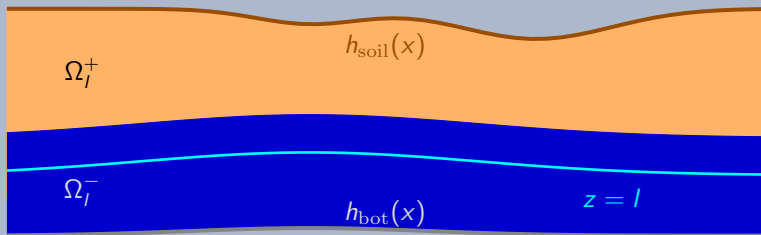
Flow described by the 3d-Richards model



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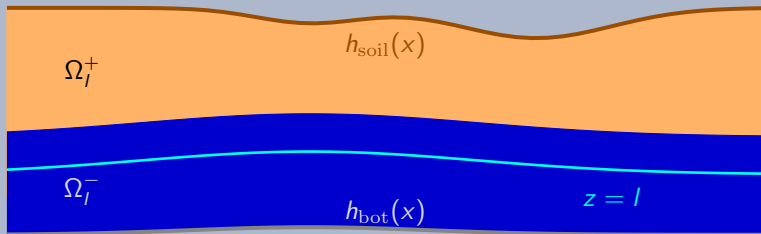
Coupled model in physical variables



Flow in short time scale

$$\frac{\partial \theta(P)}{\partial t} - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) = 0 \quad \text{in }]h_{\text{bot}}, h_{\text{soil}}[$$

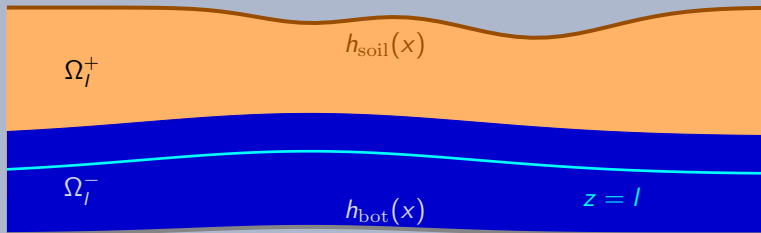
Coupled model in physical variables



Flow in short time scale

$$\frac{\partial \theta(P)}{\partial t} - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) = 0 \quad \text{in } \Omega_l^+$$
$$\underbrace{\frac{\partial \theta(P)}{\partial t}}_{=0} - \frac{\partial}{\partial z} \left(\underbrace{k(P)}_{=1} \frac{\partial H}{\partial z} \right) = 0 \quad \text{in } \Omega_l^-$$

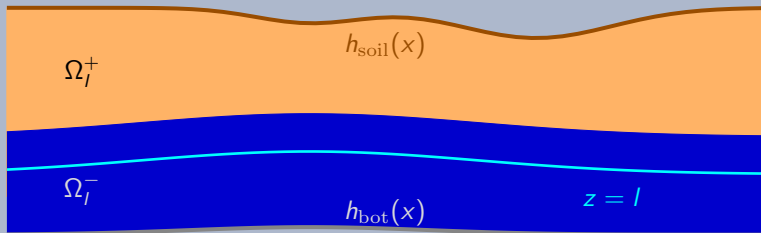
Coupled model in physical variables



Flow in short time scale

$$\frac{\partial \theta(P)}{\partial t} - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) = 0 \quad \text{in } \Omega_I^+$$
$$-\frac{\partial^2 H}{\partial z^2} = 0 \quad \text{in } \Omega_I^-$$

Coupled model in physical variables



Flow in short time scale

$$\frac{\partial \theta(P)}{\partial t} - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) = 0 \quad \text{in } \Omega_l^+$$
$$H(t, x, z) = H(t, x) \quad \text{in } \Omega_l^-$$

Coupled model in physical variables

Flow in short time scale

$$\begin{aligned} \frac{\partial \theta(P)}{\partial t} - \frac{\partial}{\partial z} \left(k(P) \frac{\partial H}{\partial z} \right) &= 0 && \text{in } \Omega_I^+ \\ H(t, x, z) &= H(t, x) && \text{in } \Omega_I^- \end{aligned}$$

Long time problem

$$\begin{cases} H(t, x, z) = H(t, x) \\ \frac{\partial}{\partial t} \int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta(P) dz - \operatorname{div}_x \left(\tilde{k}(\tilde{H}) \nabla_x H \right) = 0 \end{cases}$$

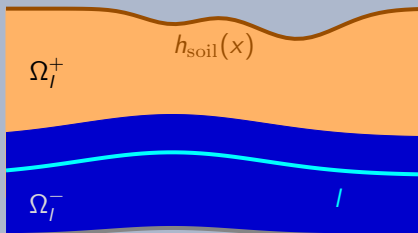
Coupled model in physical variables

Flow in Ω_l^+ : 1d-Richards

{

Flow in Ω_l^- :

$$H(t, x, z) = \tilde{H}(t, x)$$



Evolution of Hydraulic head:

Evolution of interface: 2 choices

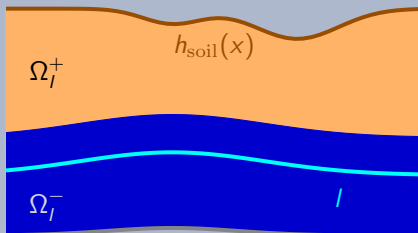
Coupled model in physical variables

Flow in Ω_I^+ : 1d-Richards

$$\begin{cases} \frac{\partial \theta(P)}{\partial t} + \frac{\partial u}{\partial z} = 0 \\ u = -k(P) \frac{\partial H}{\partial z} \\ H = \tilde{H} \end{cases} \quad \text{on } \Gamma_I$$

Flow in Ω_I^- :

$$H(t, x, z) = \tilde{H}(t, x)$$



Evolution of Hydraulic head:

Evolution of interface: 2 choices

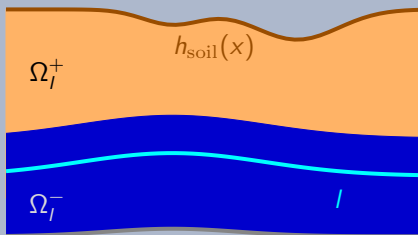
Coupled model in physical variables

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Flow in Ω_I^- :

$$H(t, x, z) = \tilde{H}(t, x)$$



Evolution of Hydraulic head:

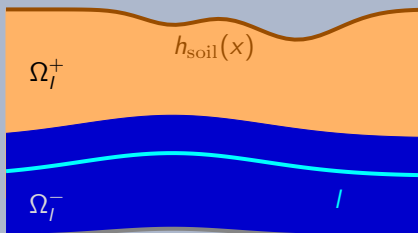
$$\frac{\partial}{\partial t} \left(\int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta(P) dz \right) - \text{div}_x \left(\tilde{k}(\tilde{H}) \nabla_x \tilde{H} \right) = 0$$

Evolution of interface: 2 choices

Coupled model in physical variables

Flow in Ω_I^+ : 1d-Richards

$$\begin{cases} \frac{\partial \theta(P)}{\partial t} + \frac{\partial u}{\partial z} = 0 \\ u = -k(P) \frac{\partial H}{\partial z} \\ H = \tilde{H} \end{cases} \quad \text{on } \Gamma_I$$



Flow in Ω_I^- :

$$H(t, x, z) = \tilde{H}(t, x)$$

Evolution of Hydraulic head:

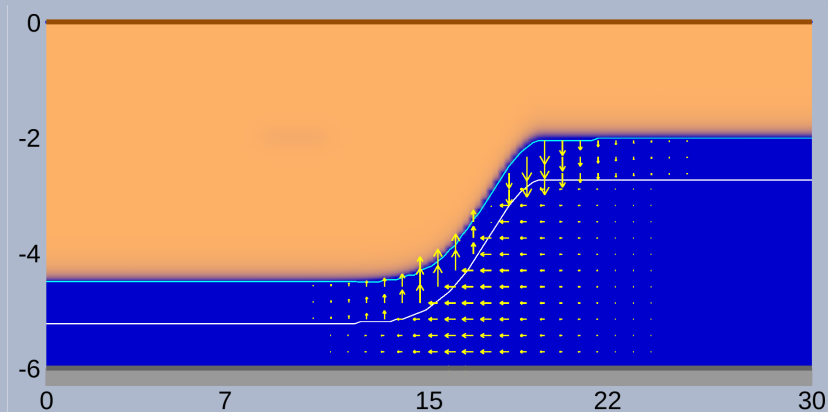
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Evolution of interface: 2 choices

- l is given a priori.
- l depends on the other variables

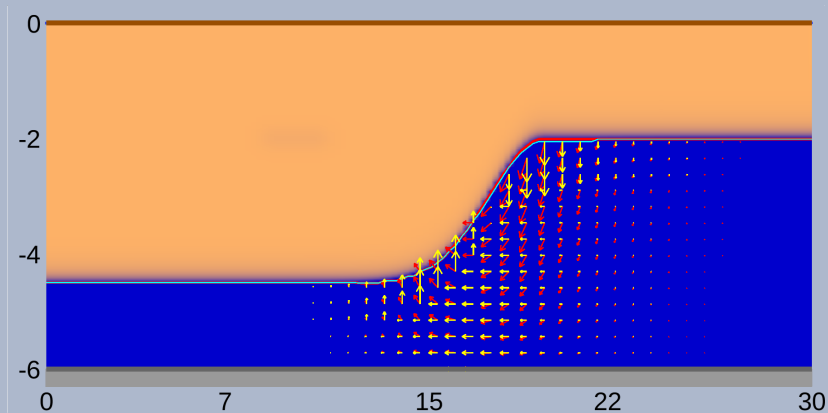
Numerical experiment

Flow described by the "Dupuit-Richards" model



Numerical experiment

Comparison between the two models



Summary of the coupled model for Richards

Geometrical coupling of dominant behavior

- Domain splitted **into too parts**
- **Upper part:** vertical Richards (fast component)
- **Lower part:** horizontal flow (slow component)

Advantages

- Coupled model and 3d-Richards models: **same effective behaviors**
 - ↔ at each time scale
- **Numerically efficient**
 - ↔ 2d problem coupled with many 1d problems
 - ↔ mass-conservative
- **Can be tuned** thanks to l

Drawbacks

- Bad description of water fluxes
 - ↔ jump from vertical to horizontal at level l .

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Implicit time scheme

$(t_n)_n$: discretization of $[0, T]$, δ_t : time step.

$$P^n(x, z) \simeq P(t_n, x, z), \quad H^n(x) \simeq H(t_n, x).$$

Evolution of Hydraulic head:

$$\frac{1}{\delta_t} \left(\int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta(P^n) dz - \int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta(P^{n-1}) dz \right) - \text{div}_x \left(\tilde{k}(\tilde{H}^n) \nabla_x \tilde{H}^n \right) = 0$$

1d-Richards

$$\begin{cases} \frac{\theta(P^n) - \theta(P^{n-1})}{\delta_t} + \frac{\partial u^n}{\partial z} = 0 & \text{in }]I^n, h_{\text{soil}}[\\ u^n = -k(P^n) \frac{\partial H^n}{\partial z} & \text{in }]I^n, h_{\text{soil}}[\\ H^n = \tilde{H}^n & \text{on } \Gamma_I^n \end{cases}$$

Implicit time scheme

$(t_n)_n$: discretization of $[0, T]$, δ_t : time step.

$$P^n(x, z) \simeq P(t_n, x, z), \quad H^n(x) \simeq H(t_n, x).$$

Evolution of Hydraulic head:

$$\frac{1}{\delta_t} \left(\int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta(P^n) dz - \int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta(P^{n-1}) dz \right) - \text{div}_x \left(\tilde{k}(\tilde{H}^n) \nabla_x \tilde{H}^n \right) = 0$$

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1d-Richards: Solved to compute Λ

$$\begin{cases} \frac{\theta(P^n) - \theta(P^{n-1})}{\delta_t} + \frac{\partial u^n}{\partial z} = 0 & \text{in }]I^n, h_{\text{soil}}[\\ u^n = -k(P^n) \frac{\partial H^n}{\partial z} & \text{in }]I^n, h_{\text{soil}}[\\ H^n = \tilde{H}^n & \text{on } \Gamma_{I^n} \end{cases}$$

Implicit time scheme

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$$\Lambda(\tilde{H}^n) - \operatorname{div}_x \left(\tilde{k}(H^n) \nabla_x \tilde{H}^n \right) = 0$$

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Picard's fix-point strategy

$$A^{n_0} = A^{n-1}, \quad A^{n_k} \rightarrow A^n \quad \text{for } A = P, H, h, u .$$

Evolution of Hydraulic head :

$$\Lambda(\tilde{H}^{n_k}) - \operatorname{div}_x \left(\tilde{k}(\tilde{H}^{n_k}) \nabla_x \tilde{H}^{n_k} \right) = 0$$

1d-Richards: used to compute Λ

$$\begin{cases} \frac{\theta(P^{n_k}) - \theta(P^{n-1})}{\delta_t} + \frac{\partial u^{n_k}}{\partial z} = 0 & \text{in }]I^n, h_{\text{soil}}[\\ u^{n_k} = -k(P^{n_k}) \frac{\partial H^{n_k}}{\partial z} & \text{in }]I^n, h_{\text{soil}}[\\ H^{n_k} = \tilde{H}^{n_k} & \text{on } \Gamma_{I^n} \end{cases}$$

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$$A^{n_0} = A^{n-1}, \quad A^{n_k} \rightarrow A^n \quad \text{for } A = P, H, h, u .$$

Evolution of Hydraulic head :

$$\underbrace{\Lambda(\tilde{H}^{n_k})}_{\simeq \Lambda(\tilde{H}^{n_{k-1}})} - \operatorname{div}_x \left(\underbrace{\tilde{k}(H^{n_k})}_{\simeq \tilde{k}(\tilde{H}^{n_{k-1}})} \nabla_x \tilde{H}^{n_k} \right) = 0$$

1d-Richards: used to compute Λ

$$\begin{cases} \frac{\theta(P^{n_k}) - \theta(P^{n-1})}{\delta_t} + \frac{\partial u^{n_k}}{\partial z} = 0 & \text{in }]I^n, h_{\text{soil}}[\\ u^{n_k} = -k(P^{n_k}) \frac{\partial H^{n_k}}{\partial z} & \text{in }]I^n, h_{\text{soil}}[\\ H^{n_k} = \tilde{H}^{n_k} & \text{on } \Gamma_{I^n} \end{cases}$$

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Evolution of Hydraulic head :

$$\Lambda(\tilde{H}^{n_{k-1}}) - \operatorname{div}_x \left(\tilde{k}(\tilde{H}^{n_{k-1}}) \nabla_x \tilde{H}^{n_k} \right) = 0$$

1d-Richards: used to compute Λ

$$\begin{cases} \frac{\theta(P^{n_k}) - \theta(P^{n-1})}{\delta_t} + \frac{\partial u^{n_k}}{\partial z} = 0 & \text{in }]I^n, h_{\text{soil}}[\\ u^{n_k} = -k(P^{n_k}) \frac{\partial H^{n_k}}{\partial z} & \text{in }]I^n, h_{\text{soil}}[\\ H^{n_k} = \tilde{H}^{n_k} & \text{on } \Gamma_{I^n} \end{cases}$$

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- **Good point:** if $(H^{n_k})_k$ converges, the limit is H^n !
-

Picard's fix-point strategy

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$$\Lambda(\tilde{H}^{n_{k-1}}) - \operatorname{div}_x \left(\tilde{k}(\tilde{H}^{n_{k-1}}) \nabla_x \tilde{H}^{n_k} \right) = 0$$

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- **Good point:** if $(H^{n_k})_k$ converges, the limit is H^n !
- **Bad point:** it does not converges...

Pseudo-Newton strategy

$$A^{n_0} = A^{n-1}, \quad A^{n_k} \rightarrow A^n \quad \text{for } A = P, H, h, u .$$

Evolution of Hydraulic head :

$$\Lambda(\tilde{H}^{n_k}) - \operatorname{div}_x \left(\tilde{k}(\tilde{H}^{n_k}) \nabla_x \tilde{H}^{n_k} \right) = 0$$

1d-Richards: used to compute Λ

$$\begin{cases} \frac{\theta(P^{n_k}) - \theta(P^{n-1})}{\delta_t} + \frac{\partial u^{n_k}}{\partial z} = 0 & \text{in }]I^n, h_{\text{soil}}[\\ u^{n_k} = -k(P^{n_k}) \frac{\partial H^{n_k}}{\partial z} & \text{in }]I^n, h_{\text{soil}}[\\ H^{n_k} = \tilde{H}^{n_k} & \text{on } \Gamma_I^n \end{cases}$$

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Evolution of Hydraulic head :

$$\underbrace{\Lambda(\tilde{H}^{n_k})}_{\simeq \Lambda(\tilde{H}^{n_{k-1}}) + \Lambda'(\tilde{H}^{n_{k-1}})(\tilde{H}^{n_k} - \tilde{H}^{n_{k-1}})} - \operatorname{div}_x \left(\underbrace{\tilde{k}(H^{n_k})}_{\simeq \tilde{k}(\tilde{H}^{n_{k-1}})} \nabla_x \tilde{H}^{n_k} \right) = 0$$

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$$\Lambda(\tilde{H}^{n_{k-1}}) + \Lambda'(\tilde{H}^{n_{k-1}})(\tilde{H}^{n_k} - \tilde{H}^{n_{k-1}}) - \operatorname{div}_x \left(\tilde{k}(\tilde{H}^{n_{k-1}}) \nabla_x \tilde{H}^{n_k} \right) = 0$$

1d-Richards: used to compute Λ

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Evolution of Hydraulic head :

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- **Good point:** if $(H^{n_k})_k$ converges, the limit is H^n !
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- **Good point:** if $(H^{n_k})_k$ converges, the limit is H^n !
- **Bad point:** none !

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- 1 Geometry and 3d-Richards equations
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- 5 Considering reactive transport**
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Chemical equilibrium: example of nitrate creation

Morel table (mobile/fixed)

	Primary Species			Secondary Species				
	H^+	NH_4^+	$SoilH$	OH^-	NH_3	$Soil^-$	$SoilH_2^+$	$SoilNH_4$
H^+	1			-1	-1	-1	1	-1
NH_4^+		1			1			1
$SoilH$			1			1	1	1
K				10^{-9}	10^{-14}	10^{-7}	10^2	0.4

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K				10^{-9}	10^{-14}	10^{-7}	10^2	0.4

Unknwnowns

- **Concentration** of species

$[H^+]$, $[NH_4^+]$, $[SoilH]$ (Primary)

$[OH^-]$, $[NH_3]$, $[Soil^-]$, $[SoilH_2^+]$, $[SoilNH_4]$ (Secondary)



Chemical equilibrium: example of nitrate creation

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Unknwnowns

- **Concentration** of species

$[H^+]$, $[NH_4^+]$, $[SoilH]$ (Primary)

$[OH^-]$, $[NH_3]$, $[Soil^-]$, $[SoilH_2^+]$, $[SoilNH_4]$ (Secondary)

- Total of each **component**: T_{H^+} , $T_{NH_4^+}$, T_{SoilH}

Chemical equilibrium: example of nitrate creation

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Unknwnowns

- **Concentration** of species
 $[H^+]$, $[NH_4^+]$, $[SoilH]$ (Primary)
 $[OH^-]$, $[NH_3]$, $[Soil^-]$, $[SoilH_2^+]$, $[SoilNH_4]$ (Secondary)
- Total of each **component**: T_{H^+} , $T_{NH_4^+}$, T_{SoilH}
- **mobile** and **fixed** contrib.:
 $T_{H^+}^m$, $T_{NH_4^+}^m$, T_{SoilH}^m ,
 $T_{H^+}^f$, $T_{NH_4^+}^f$, T_{SoilH}^f

Chemical equilibrium: example of nitrate creation

Morel table (mobile/fixed)

	Primary Species			Secondary Species				
	H^+	NH_4^+	$SoilH$	OH^-	NH_3	$Soil^-$	$SoilH_2^+$	$SoilNH_4$
H^+	1			-1	-1	-1	1	-1
NH_4^+		1			1			1
$SoilH$			1			1	1	1
K				10^{-9}	10^{-14}	10^{-7}	10^2	0.4

Unknwnowns

- **Concentration** of species

$$a = (a^m, a^f) = ([H^+], [NH_4^+], [SoilH]) \dots \dots \dots \text{(Primary)}$$

$$b = (b^m, b^f) = ([OH^-], [NH_3], [Soil^-], [SoilH_2^+], [SoilNH_4]) \dots \dots \text{(Sec.)}$$

- Total of each **component**: $T = (T_{H^+}, T_{NH_4^+}, T_{SoilH})$

$$T^m = (T_{H^+}^m, T_{NH_4^+}^m, T_{SoilH}^m)$$

- **mobile** and **fixed** contrib.:

$$T^f = (T_{H^+}^f, T_{NH_4^+}^f, T_{SoilH}^f)$$

Chemical Equilibrium: example of nitrate creation

$$\text{Morel Table} \implies \begin{cases} \text{Stoichiometry matrix } \begin{pmatrix} \mathbf{I}_d & \mathbf{M} \end{pmatrix} \\ \mathbf{K} = (K_j)_j = (1, 10^{-14}, 1, 10^{-9}, 1, 10^{-7}, 10^2, 0.4) \end{cases}$$

with

- **Linear relations**

$$\mathbf{T} = \mathbf{a} + \mathbf{M}\mathbf{b}, \quad \mathbf{T} = \mathbf{T}^m + \mathbf{T}^f, \quad \mathbf{T}^m = \begin{pmatrix} \mathbf{a}^m + \mathbf{M}^m \mathbf{b}^m \\ 0 \end{pmatrix}$$

- **Non-linear relations**

$$b_i = K_i \prod_j a_j^{M_{ij}}, \quad \text{for all } i$$

i.e. $\log(K^i) = b_i - \sum_j M_{ij} \log(a_j)$ for all i

i.e.

$$\log(\mathbf{K}) = \mathbf{b} - \mathbf{M} \log(\mathbf{a})$$

Chemical Kinetic: example of nitrate creation

New Unknowns

- $[NO_3^-]$, $[O_2]$, $[Bac]$

Example of a kinetic reactions



with the reaction rate (non linear)

$$r_1 = r_1([Bac], [O_2], [NH_4^+]) = \mu[Bac] \frac{k_1}{k_1 + [Bac]} \frac{[NH_4^+]}{k_2 + [NH_4^+]} \frac{[O_2]}{k_3 + [O_2]}$$

Cinetic reactions

Acts as **non-linear** source terms in mass conservation equation.

Chemical Kinetic: example of nitrate creation

New Unknowns

- $s = ([NO_3^-], [O_2], [Bac])$

Example of a kinetic reactions



with the reaction rate (non linear)

$$r_1 = r_1([Bac], [O_2], [NH_4^+]) = \mu[Bac] \frac{k_1}{k_1 + [Bac]} \frac{[NH_4^+]}{k_2 + [NH_4^+]} \frac{[O_2]}{k_3 + [O_2]}$$

Cinetic reactions

Acts as **non-linear** source terms in mass conservation equation.

Mass-conservation equation

Notations

- $\mathbf{c} = (\mathbf{T}, \mathbf{s})$: quantities transported
- $\mathbf{e} = (\mathbf{T}^m, \mathbf{s})$: quantities transported (mobile contributions)
- $\mathbf{d} = (\mathbf{a}, \mathbf{b}, \mathbf{s})$: concentrations of every chemical species involved
- $\mathbf{r} = \mathbf{r}(\mathbf{d})$ reaction rate (one component for each cinetic reaction)

Equilibrium equations

- Linear relations

$$\mathbf{c} = \mathbf{A}\mathbf{d}, \quad \mathbf{e} = \mathbf{E}\mathbf{d}$$

- Non-linear relations

$$\log(\mathbf{K}) = \Lambda \log(\mathbf{d})$$

Mass-concervation equation

For the flow

- θ : Water saturation
- H : Hydraulic head
- \mathbf{v}^f : Fluid velocity

$$\left\{ \begin{array}{l} \frac{\partial \theta}{\partial t} + \operatorname{div} \mathbf{v}^f = g^f \\ \mathbf{v}^f = -k_r \mathbf{K}_0 \nabla H \\ + \text{limit cond.} \end{array} \right.$$

Transport of multiple reacting species

- c^i : quantity transported (in the water)
- e^i : quantity transported, mobile contrib.
- \mathbf{v}^{e^i} : velocity of the specy (diffusion + dispersion + transport)
- $\mathbf{D} = \mathbf{D}(\mathbf{v}^f)$: diffusion + dispersion tensor

$$\left\{ \begin{array}{l} \frac{\partial}{\partial t}(\theta c^i) + \operatorname{div}(\mathbf{v}^{e^i}) = g_i^c + R^i(\mathbf{d}) \\ \mathbf{v}^{e^i} := e^i \mathbf{v}^f - \theta \mathbf{D}(e^i) \nabla e^i \end{array} \right.$$

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Dominant component for the transport problem

Reactive transport

$$\begin{cases} \frac{\partial}{\partial t}(\theta c^i) + \operatorname{div}(\mathbf{v}^{e^i}) = g^{c^i} + R^i(\mathbf{d}) \\ \mathbf{v}^{e^i} := e^i \mathbf{v}^f - \theta D \nabla e^i \end{cases}$$

Chemical equilibrium

$$\begin{cases} \mathbf{c} = \mathbf{A} \mathbf{d}, \\ \mathbf{e} = \mathbf{A}^m \mathbf{d}, \\ \log \mathbf{K} = \mathbf{B} \log(\mathbf{d}) \end{cases}$$

Transport in short time scale: 1d-vertical

$$\begin{cases} \frac{\partial}{\partial t}(\theta c^i) + \frac{\partial v_3^{e^i}}{\partial z} = g^{c^i} + R^i(\mathbf{d}) \\ v_3^{e^i} := e^i v_3^f - \theta D \frac{\partial e^i}{\partial z} \end{cases}$$

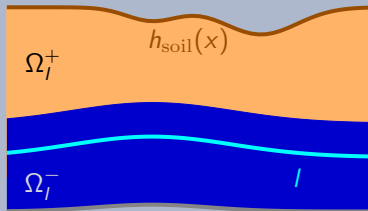
Long time problem: 2d-horizontal

$$\begin{cases} d^i(t, x, z) = \tilde{d}^i(t, x) \\ \frac{\partial}{\partial t} \int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta c^i dz - \operatorname{div}_x \tilde{\mathbf{w}}^i = 0 \\ \tilde{\mathbf{w}}^i = \tilde{e}^i \tilde{\mathbf{w}}^f - \tilde{\theta} D^i \nabla_{\bar{x}} \tilde{e}^i \end{cases}$$

Coupled model in physical variables

1d-vertical transport in Ω_l^+

$$\begin{cases} \frac{\partial(\theta c^i)}{\partial t} + \frac{\partial v_3^{e^i}}{\partial z} = g^{c^i} + R^i(\mathbf{d}) \\ v_3^{e^i} := e^i v_3^f - \theta D \frac{\partial e^i}{\partial z} \\ c^i = \tilde{c}^i \quad \text{for } z = l \end{cases}$$



Constant concentrations in Ω_l^- :

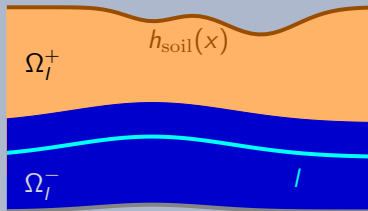
$$\mathbf{d}(t, x, z) = \tilde{\mathbf{d}}(t, x), \quad c(t, x, z) = \tilde{c}(t, x), \quad \mathbf{e}(t, x, z) = \tilde{\mathbf{e}}(t, x)$$

2d problem

Coupled model in physical variables

1d-vertical transport in Ω_l^+

$$\begin{cases} \frac{\partial(\theta c^i)}{\partial t} + \frac{\partial v_3^{e^i}}{\partial z} = g^{c^i} + R^i(\mathbf{d}) \\ v_3^{e^i} := e^i v_3^f - \theta D \frac{\partial e^i}{\partial z} \\ c^i = \tilde{c}^i \quad \text{for } z = l \end{cases}$$



Constant concentrations in Ω_l^- :

$$\mathbf{d}(t, x, z) = \tilde{\mathbf{d}}(t, x), \quad c(t, x, z) = \tilde{c}(t, x), \quad \mathbf{e}(t, x, z) = \tilde{\mathbf{e}}(t, x)$$

2d problem

$$\frac{\partial}{\partial t} \int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta c^i dz - \text{div}_x \tilde{\mathbf{w}}^i = \int_{h_{\text{bot}}}^{h_{\text{soil}}} R^i(\mathbf{d}) dz, \quad \tilde{\mathbf{w}}^i = \tilde{e}^i \tilde{\mathbf{w}}^f - \tilde{\theta} D^i \nabla_{\bar{x}} \tilde{e}^i$$

Full Coupled reactive-transport problem

Coupled Reactive transport

$$\frac{\partial}{\partial t}(\theta c^i) + \frac{\partial v_3^{e^i}}{\partial z} = R^i(\mathbf{d}) \quad \text{in } \Omega_I^+$$
$$\frac{\partial}{\partial t} \left(\int_{h_{\text{bot}}}^{h_{\text{soil}}} \theta c^i dz \right) - \text{div}_x \tilde{\mathbf{w}}^i = \int_{h_{\text{bot}}}^{h_{\text{soil}}} R^i(\mathbf{d}) dz,$$
$$v_3^{e^i} = e^i v_3^f - \theta D \frac{\partial e^i}{\partial z} \quad \text{in } \Omega_I^+$$
$$\tilde{\mathbf{w}}^i = \tilde{e}^i \tilde{\mathbf{w}}^f - \tilde{\theta} D^i \nabla_{\bar{x}} \tilde{e}^i$$

In Ω_I^-

$$\mathbf{d}(t, x, z) = \tilde{\mathbf{d}}(t, x), \quad \mathbf{c}(t, x, z) = \tilde{\mathbf{c}}(t, x), \quad \mathbf{e}(t, x, z) = \tilde{\mathbf{e}}(t, x)$$

Chemical equilibrium:

$$\mathbf{c} = \mathbf{A} \mathbf{d}, \quad \mathbf{e} = \mathbf{A}^m \mathbf{d}, \quad \log \mathbf{K} = \mathbf{B} \log(\mathbf{d})$$

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+2d problem

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Plan

- 1 Geometry and 3d-Richards equations
- 2 Dominant components of the flow: effective problems
- 3 A coupled model coupling dominant components
- 4 Numerical scheme
- 5 Considering reactive transport
- 6 Reactive transport: Dominant component and coupled model**
- 7 Numerical scheme

Implicit time scheme

$(t_n)_n$: discretization of $[0, T]$, δ_t : time step.

$$\mathbf{c}^n(x, z) \simeq \mathbf{c}(t_n, x, z), \quad c^{i,n}(x, z) \simeq c^i(t_n, x, z) \dots$$

Coupled Reactive transport

$$\frac{1}{\partial t}(\theta^n c^{i,n} - \theta^{n-1} c^{i,n-1}) + \frac{\partial v_3^{e^{i,n}}}{\partial z} = R^i(\mathbf{d}^n) \quad \text{in } \Omega_I^+$$

$$v_3^{e^{i,n}} = e^{i,n} v_3^{f,n} - \theta D \frac{\partial e^{i,n}}{\partial z} \quad \text{in } \Omega_I^+$$

In Ω_I^-

$$\mathbf{d}^n(t, x, z) = \tilde{\mathbf{d}}^n(t, x), \quad \mathbf{c}^n(t, x, z) = \tilde{\mathbf{c}}^n(t, x), \quad \mathbf{e}^n(t, x, z) = \tilde{\mathbf{e}}^n(t, x)$$

+2d problem

Chemical equilibrium:

$$\mathbf{c}^n = \mathbf{A} \mathbf{d}^n, \quad \mathbf{e}^n = \mathbf{A}^m \mathbf{d}^n, \quad \log K = \mathbf{B} \log(\mathbf{d}^n)$$

Picard fixed-point

Strategy

- At each time scale n , we know: \mathbf{d}^{n-1}
- We build a sequence $(\mathbf{d}^{n_k})_k$ such that

$$\mathbf{d}^{n_0} = \mathbf{d}^{n-1}, \quad \mathbf{d}^{n_k} \rightarrow \mathbf{d}^n: \text{ solution of the implicit problem}$$

Coupled Reactive transport: give the solution c^{i,n_k}

We know $\mathbf{c}^{n_{k-1}}$, $\mathbf{e}^{n_{k-1}}$, $\mathbf{d}^{n_{k-1}}$, $\tilde{\mathbf{c}}^{n_{k-1}}$...

$$\frac{1}{\partial t}(\theta^n c^{i,n_k} - \theta^{n-1} c^{i,n-1}) + \frac{\partial \hat{v}_3^{c^{i,n_k}}}{\partial z} = R^i(\mathbf{d}^{n_{k-1}}) \quad \text{in } \Omega_I^+$$

$$\hat{v}_3^{c^{i,n_k}} = v_3^{c^{i,n_k}} + (v_3^{e^{i,n_{k-1}}} - v_3^{c^{i,n_{k-1}}}) \quad \text{in } \Omega_I^+$$

In Ω_I^-

$$\mathbf{d}^n(t, x, z) = \tilde{\mathbf{d}}^n(t, x), \quad \mathbf{c}^n(t, x, z) = \tilde{\mathbf{c}}^n(t, x), \quad \mathbf{e}^n(t, x, z) = \tilde{\mathbf{e}}^n(t, x)$$

+2d problem

Chemical equilibrium: give \mathbf{d}^{n_k} and \mathbf{e}^{n_k}

$$\mathbf{c}^{n_k} = \mathbf{A} \mathbf{d}^{n_k}, \quad \mathbf{e}^{n_k} = \mathbf{A}^m \mathbf{d}^{n_k}, \quad \log K = \mathbf{B} \log(\mathbf{d}^{n_k})$$

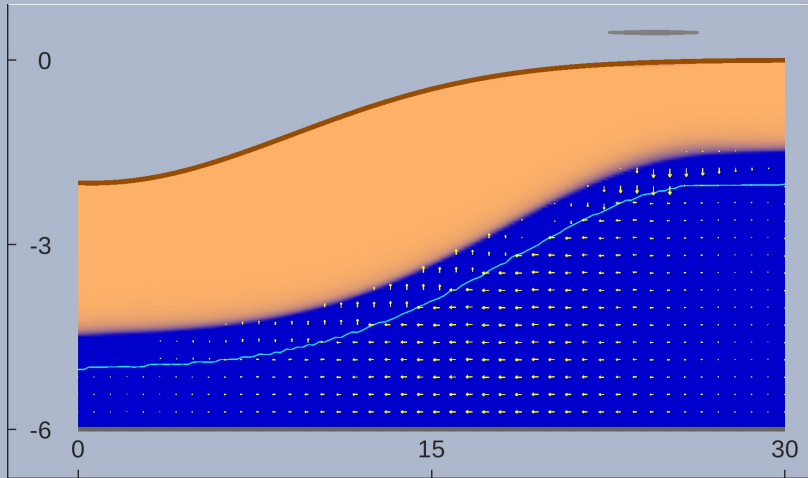
Convergence of Picard algorithm

Difficulty: During resolution c^{n_k} must be positive !

- Need a transport scheme **preserving positivity**
- Can be false **before** the convergence being reached...
 - ↪ we use a truncature procedure to obtain convergence

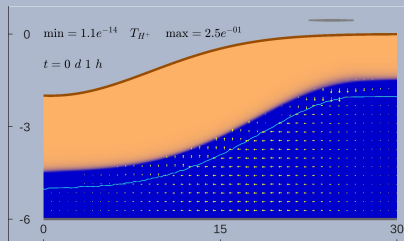
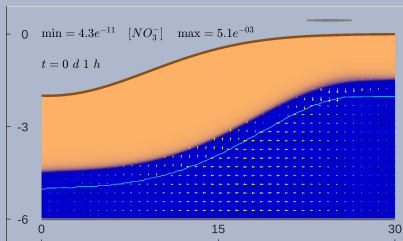
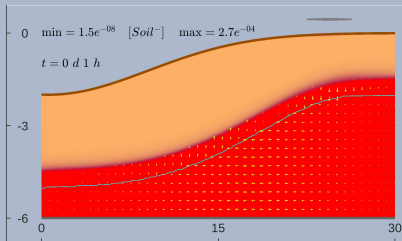
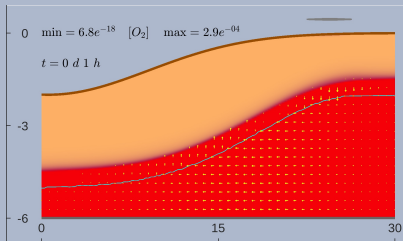
First transport experiment

One specy, **only transport** (no cinetic, no equilibrium)



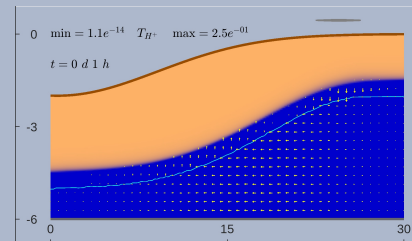
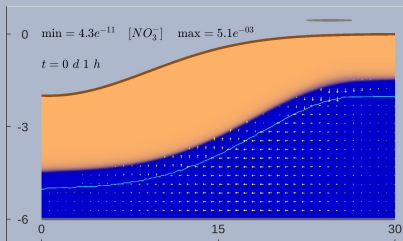
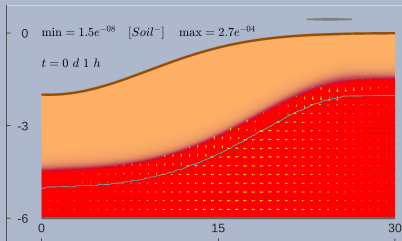
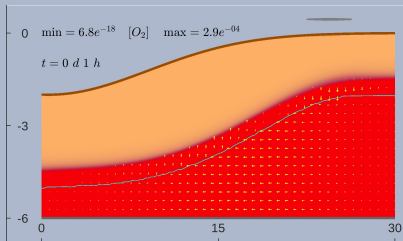
Complete experiment

11 species, 6 transported quantities, equilibrium and kinetic

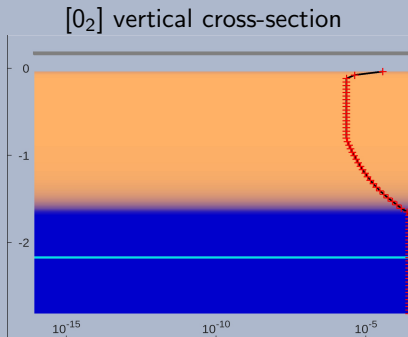
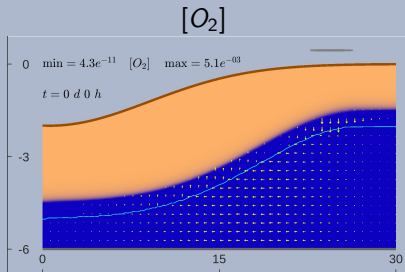


Complete experiment

11 species, 6 transported quantities, equilibrium and kinetic



Complete experiment: cross-section view



Complete experiment: cross-section view

